

1. Introduction

- Bubble growth in rhyolitic melts is a primary control on some of the largest explosive eruptions. However, vesiculation remains poorly constrained.
- Few studies have captured in-situ vesiculation of a rhyolitic melt, but have rather relied on interpretation of quenched natural or experimental samples. The previous in-situ study^[1] followed vesiculation in water-poor (~0.14 wt%) rhyolitic melt at P = 1 At.
- This work aims to provide measurements of rapid in-situ vesiculation in more water-rich rhyolitic melt from high resolution imagery. Results are relevant to post-fragmentation magma vesiculation.

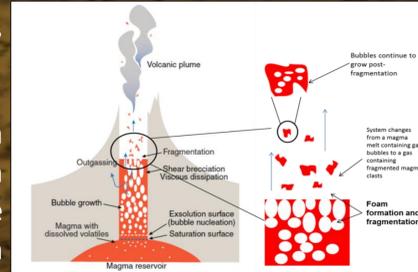


Fig 1. The stages of bubble nucleation and growth leading to magma fragmentation in a volcanic conduit. Bubbles continue to grow at atmospheric pressure post-fragmentation. Modified from [2].

2. Methods

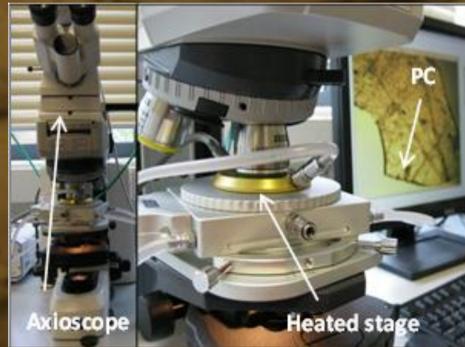


Fig 2. Hot-stage microscope setup³. Sample is placed inside a ceramic furnace (Linkam TS1500 heated stage), mounted on a Zeiss AxioScope.

- Using the technique of Applegarth et al (2013), thin wafers (~ 100 μm thick) of obsidian (0.97 wt% H₂O) from the 2008 eruption at Chaitén, Chile were held from 5 minutes up to 2 days in the hotstage at between 575 °C and 875 °C. All experiments were conducted at 1 atm and therefore do not consider growth by decompression.
- The in-situ growth of many individual bubbles were recorded directly to PC and then measured using particle tracking code written in MATLAB.

3. Gas loss from sample surface

- The potential for sample dehydration was considered by estimating the extent of diffusive degassing from wafer surfaces using simple diffusion models^[4]
- Dehydration was found to be negligible during brief high temperature experiments but became increasingly important for slower, lower-temperature experiments

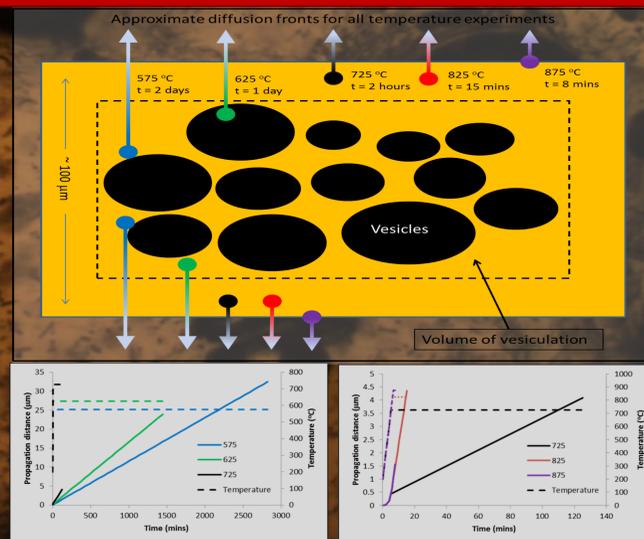


Fig 3. Diffusive degassing from wafer surface during heating experiments

4. The physical processes of bubble growth

- Five stages of bubble growth were directly observed (Fig. 4).
- Most rapid average bubble growth rate at 875 °C (1.27 μm s⁻¹; η = 10^{7.1} Pa s)
- Slowest bubble growth rate at 725 °C (0.02 μm s⁻¹; η = 10^{9.20} Pa s)
- No bubble growth was noticeable below 725 °C.
- Growth rates decreased with time, as reported in [1]. However, growth rate decreases observed here are due to bubble-bubble interactions.

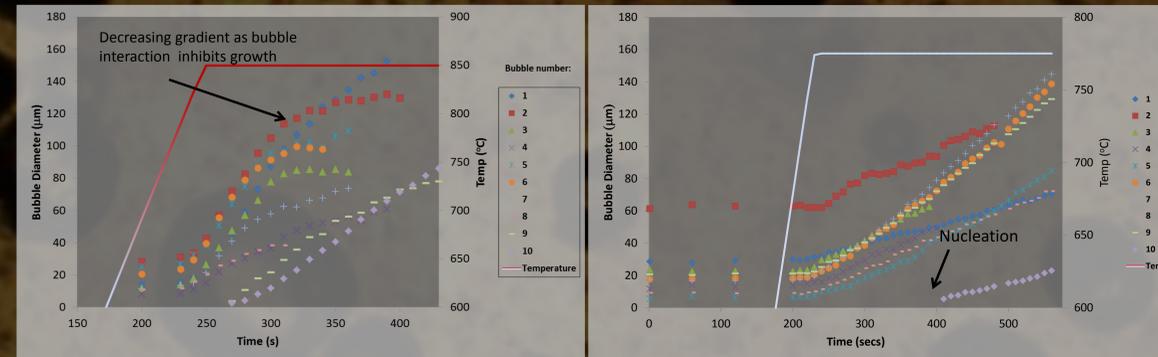


Fig 5. Bubble growth over a period of > 200 seconds at temperatures of 850 °C and 775 °C.

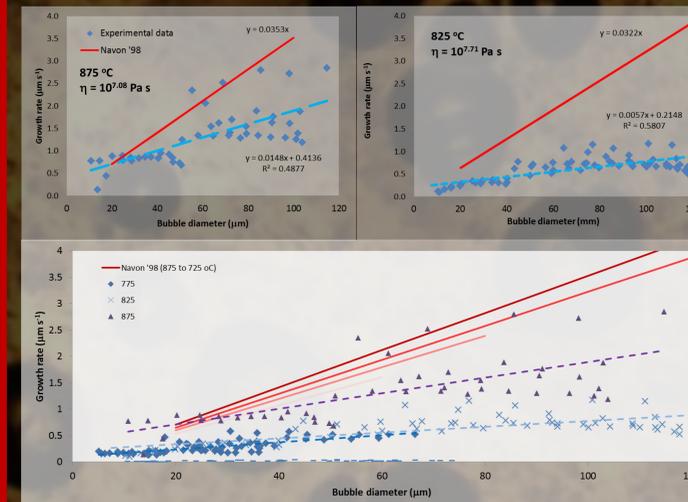


Fig 6. Comparison of experimental bubble growth rates compared to those modelled by Navon et al, 1998.

5. Further findings and conclusions

- Water content strongly influences vesicle growth rates, which are ~7 times higher in the water-rich Chaitén rhyolite than the GOVC peralkaline rhyolite (0.14 wt%) used in [1].
- We estimate bubble nucleation rate (J) of >1.5 x 10¹⁰ m⁻³ s⁻¹, from change in bubble number through time. This matches the lower end of J values from decompression experiments^[6]. High nucleation rates occurred for ~30 seconds, prior to bubble number reduction due to coalescence during foaming.

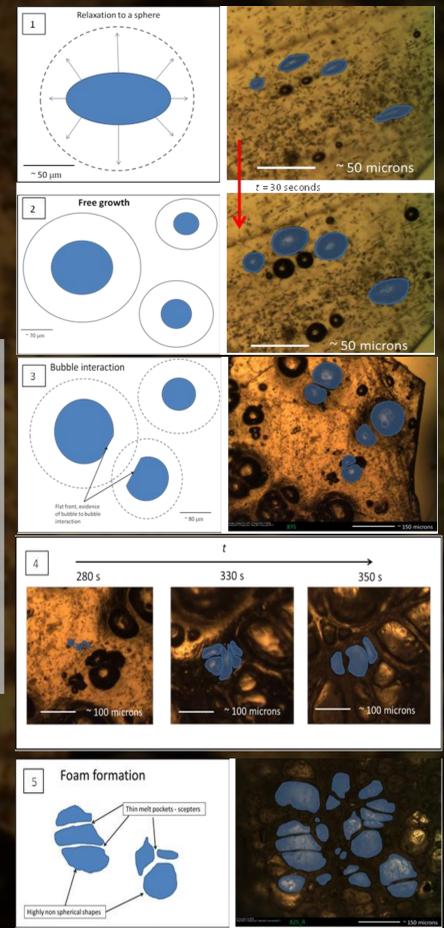


Fig 4. Five stages of bubble growth

- Experimentally obtained bubble growth rates were compared to the predicted growth rates modelled by Navon et al (1998), Fig. 6, using: $P_g - P_f = \frac{2\sigma}{R} + 4\eta \frac{V_r}{R}$
- Where the bubble overpressure (P_g) - ambient pressure (P_f) is assumed to be initially constant, and the effect of surface tension (σ) is ignored for bubble radius > 5.5 μm ∴ Growth rate (V_r) = R/4η
- At all temperatures modelled growth rates exceed those determined experimentally; there is a closer fit at higher temperatures.
- This finding is contrary to [1], whose modelled growth rate fitted better at lower temperatures. This discrepancy may relate to the varied water contents of samples studied.

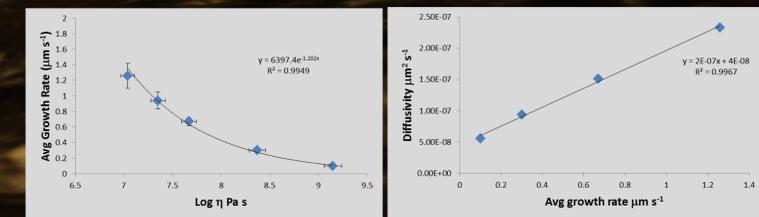


Fig 7. Bubble growth rate (V_r) increases with decreasing melt viscosity(η) where: V_r ≈ exp(-1.202η). At higher temperatures (> 725 °C) there is a positive linear relationship between bubble growth and modelled diffusivity.

References
¹Bagdassarov, N.S et al (1996) Rhyolite magma degassing: an experimental study of melt vesiculation. *Bulletin of volcanology*, 57, 587-601
²Gonnermann, H.M and Manga, M (2007) The fluid mechanics inside a volcano. *Annual review of fluid mechanics*, 39, 321-356
³Applegarth, L.J., Tuffen, H., James, M.R., Pinkerton, H., Cashman, K.V. (2013) Direct observations of degassing-induced crystallization in basaltic. *Geology* doi:10.1130/G33641.1. See poster V11C-2780
⁴Zhang, Y (1999) H₂O in rhyolitic glasses and melts: Measurements, speciation, solubility, and diffusion. *Reviews in Geophysics*.
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⁶Ciuzel N. et al. (2008) *Contrib Mineral Petrol* (2008) 156:745-763; Gardner & Ketcham (2011) *Contrib Mineral Petrol* 162:929-943.